

The perovskite to post-perovskite transition in CaIrO_3 : Clapeyron slope and changes in bulk and shear moduli by density functional theory

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Abstract

The thermodynamic properties of perovskite (pv) and post-perovskite (ppv) of CaIrO_3 are derived from total energy calculations using density functional theory. Negative molar volume and enthalpy changes of $0.40 \text{ cm}^3 \text{ mol}^{-1}$ and 11.8 kJ mol^{-1} for the pv to ppv transition at 0 K stabilize ppv at low temperatures and high-pressures. Vibrational entropies calculated in the harmonic approximation, using the direct method, favour pv with increasing temperatures ($105.5 \text{ J K}^{-1} \text{ mol}^{-1}$ for pv versus $99.2 \text{ J K}^{-1} \text{ mol}^{-1}$ for ppv at 298 K). A main reason for the lower entropy of post-perovskite compared to perovskite is probably related to constraints on certain vibrational modes imposed by edge-sharing of octahedra in post-perovskite. The Clapeyron slope of the pv–ppv phase boundary is deduced from the calculated enthalpy and volume of transition in conjunction with an experimental transition temperature. The resulting dp/dT of 18 MPa K^{-1} is in good agreement with experimental determinations reported in literature. The high-pressure properties were calculated from the variation of the total energy with volume using the Murnaghan (M) and the Birch–Murnaghan (B–M) equations-of-state. The two methods give the same values for the bulk modulus but somewhat different values for the pressure derivative of the bulk modulus: $K_0 = 178 \text{ GPa}$ and $K'_0 = 2.8$ (M) and 3.3 (B–M) for pv compared to $K_0 = 164 \text{ GPa}$ and $K'_0 = 3.9$ (M) and 4.0 (B–M) for ppv. By holding K'_0 for pv fixed at 4.0, K_0 is reduced to 172 GPa. The bulk moduli at zero pressure were also derived through calculation of the elastic constants giving $K_0 = 172 \text{ GPa}$ for pv compared to $K_0 = 157 \text{ GPa}$ for ppv. In order to compare the changes in the bulk and shear moduli across the pv–ppv phase transition, also the shear moduli were derived from the elastic constants. Whereas the bulk modulus decreases by about 9% from perovskite to post-perovskite, the shear modulus increases by about 15% across the transition. The elastic parameters obtained for the CaIrO_3 polymorphs are consistent with those of MgSiO_3 and with seismic velocity variations in the lowermost mantle. Our DFT-results demonstrate that CaIrO_3 may be a useful low-pressure analogue for studies of the general properties of the pv–ppv-transition relevant to the D'' layer since the negative and positive changes in bulk and shear moduli, respectively, across the pv- to ppv-transition boundary appear to be strongly linked to the nature of the two crystal structures. © 2007 Elsevier B.V. All rights reserved.

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1. Introduction

The recent discovery of a phase transition from MgSiO_3 -dominated perovskite (pv) (space group $Pbnm$) to a post-perovskite (ppv) phase with CaIrO_3 -type struc-

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ture (space group *Cmcm*) at pT -conditions corresponding to the D'' -zone (e.g. Murakami et al., 2004, 2005; Hirose et al., 2005; Mao et al., 2004) has important implications for the core to mantle heat flux and deep Earth dynamics. The physiochemical properties of coexisting perovskite and post-perovskite and the Clapeyron slope of the phase transition are critical parameters in geodynamical models. The transition from perovskite to post-perovskite is associated with a large decrease in entropy combined with a small decrease in volume, resulting in a large positive Clapeyron slope. The dp/dT -slope of the phase boundary in MgSiO_3 cannot be determined accurately in laser-heated diamond anvil cell experiments. Experiments using suitable analogue compounds where the pv – ppv -transition can be studied within the pressure ranges of the piston cylinder or multianvil apparatus may therefore be useful.

At atmospheric pressure the stable form of CaIrO_3 is the orthorhombic post-perovskite (*Cmcm*), whereas the orthorhombic perovskite (*Pbnm*) may form metastably (McDaniel and Schneider, 1972). The latter phase is stable at high temperatures at somewhat higher pressures. Hirose and Fujita (2005) and Trønnes et al. (2006) investigated the pv – ppv -transition for CaIrO_3 at 1–3 GPa and 1350–1550 °C, and found dp/dT -slopes of about 16 and 24 MPa K^{-1} , respectively. In a calorimetric study of the CaIrO_3 polymorphs, Kojitani et al. (2007) determined an even larger transition slope of about 40 MPa K^{-1} . These slopes are considerably larger than those estimated for MgSiO_3 from first principles and *ab initio* models (Iitaka et al., 2004; Oganov and Ono, 2004) and from modelling of the topography of the upper and lower D'' -discontinuities (Sidorin et al., 1999; Nakagawa and Tackley, 2004; Hernlund et al., 2004) assuming that these discontinuities are caused by the pv – ppv -transition. Clapeyron slopes ranging from 6 to 10 MPa K^{-1} have been used in such geodynamic modelling. Based on an ambient pressure thermal expansion study of ppv - CaIrO_3 , Lindsay-Scott et al. (2007) found considerably larger elastic anisotropy in CaIrO_3 compared to MgSiO_3 . This led them to conclude that the physical properties at atmospheric pressure make ppv - CaIrO_3 a very imperfect analogue for ppv - MgSiO_3 . At slightly elevated pressures, however, this discrepancy may become smaller.

The pv – ppv -transition in MgSiO_3 is unusual in that the bulk and shear moduli may change in opposite directions across the phase boundary. Several seismic discontinuities that have been observed within the D'' -zone show large lateral depth-variation and complexity (e.g. Lay et al., 2006). An uppermost discontinuity characterized by a shear wave velocity drop may be caused

by a density jump due to compositional changes. The pv – ppv -transition seems to be related to a considerable shear wave speed increase. A locally developed and generally weaker shear wave decrease closer to the core–mantle boundary can be ascribed to the reverse transition from ppv to pv caused by the large thermal gradient near the core. As pointed out by Wookey et al. (2005) based on *ab initio* DFT calculations of elastic constants, the pv - to ppv -transition in MgSiO_3 is expected to cause discontinuities involving increasing shear wave velocity, v_s , decreasing bulk sound velocity, v_ϕ , and possibly slightly decreasing or constant pressure-wave velocity, v_p . These changes are caused by a small increase in density combined with higher shear modulus and lower bulk modulus in ppv relative to pv . The bulk modulus decrease from pv to ppv has been confirmed by experiments (Ono et al., 2006; Shieh et al., 2006).

Although the CaIrO_3 and MgSiO_3 are chemically very different, the fact that pv (*Pbnm*) and ppv (*Cmcm*) have identical space groups in both compounds makes the CaIrO_3 -system an interesting study object, e.g. by experiments in the pressure and temperature range of 1–3 GPa and 1400–1550 °C. Several studies of experimental phase equilibria, single-crystal XRD characterization of bulk modulus and thermal expansivity, calorimetry and deformation mechanisms based on samples synthesized in this p,T -range are in progress or have recently been published. Here we present a density functional theory study of perovskite and post-perovskite CaIrO_3 with focus on the thermochemistry and thermophysics of the two polymorphs. Total energy calculations are used to derive the enthalpies, volumes and bulk moduli. The vibrational properties of the ground state structures are subsequently derived in the harmonic approximation using the direct method and the entropy of pv - and ppv - CaIrO_3 estimated from the vibrational densities of state.

The derived Clapeyron-slope of the ppv - to pv -transition may provide useful supplementary information on the geodynamics of D'' region. The complex seismicity of the lowermost mantle (e.g. Hernlund et al., 2004; Lay et al., 2006; Wookey et al., 2005), including the significant seismic anisotropy, is discussed in light of the calculated elastic constants and bulk and shear moduli.

2. Computational details

2.1. Structure of post-perovskite and perovskite CaIrO_3

Both structures are described in literature. However, while a detailed structural description is found for the

post-perovskite (Rodi and Babel, 1965), only lattice vectors are reported for pv-CaIrO₃ (Hirose and Fujita, 2005). For the post-perovskite, the standard *Cmcm* setting of space groups 63 is used in the present study. Ca occupies a *4c* site with fractional coordinates (0, *y*, 1/4), Ir a *4a* site (0, 0, 0), and O both a *4c* (0, *y*, 1/4) and a *8f* (0, *y*, *z*) site. For the perovskite, the non-standard *Pbnm* setting of space group 62 is used since this has been used widely in previous studies of pv- to ppv-transitions. Ca occupies a *4c* site with fractional coordinates (*x*, *y*, 1/4), Ir a *4b* site (1/2, 0, 0) and O both a *4c* (*x*, *y*, 1/4) and a *8d* (*x*, *y*, *z*) site. Polyhedron representations of the crystal structures are given in Fig. 1. While, pv-CaIrO₃ forms a three-dimensional structure based on corner-sharing octahedra only, ppv-CaIrO₃ contains a mixture of corner- and edge-sharing octahedra resulting in a layered structure. While the octahedra share corners along the *c*-axis, they share edges along the *a*-axis and the two-dimensional slabs are thus stacked along the *b*-axis.

No experimental information on the magnetic structures has to our knowledge been reported.

2.2. Structural optimization by total energy calculations

Structural optimizations with respect to all unit cell dimensions and atomic coordinates of both structures were performed using density functional theory within the generalized gradient approximation, using the PBE functional (Perdew et al., 1996) for the exchange-correlation contribution to the total energy. A sufficiently large basis of projected augmented plane waves was used as implemented in the Vienna *ab initio* simulation program (VASP) (Kresse and Hafner, 1993; Kresse and Furthmüller, 1999). Both polymorphs are here represented by periodic unit cells containing four formula units. The post-perovskite structure can alternatively be represented by a two formula unit primitive cell. To ensure high accuracy, the *k*-point density and the plane wave cutoff energy were increased until convergence. The cutoff energy for the plane wave basis was 700 eV, while the *k*-point densities were $9 \times 5 \times 5$ for *Cmcm* and

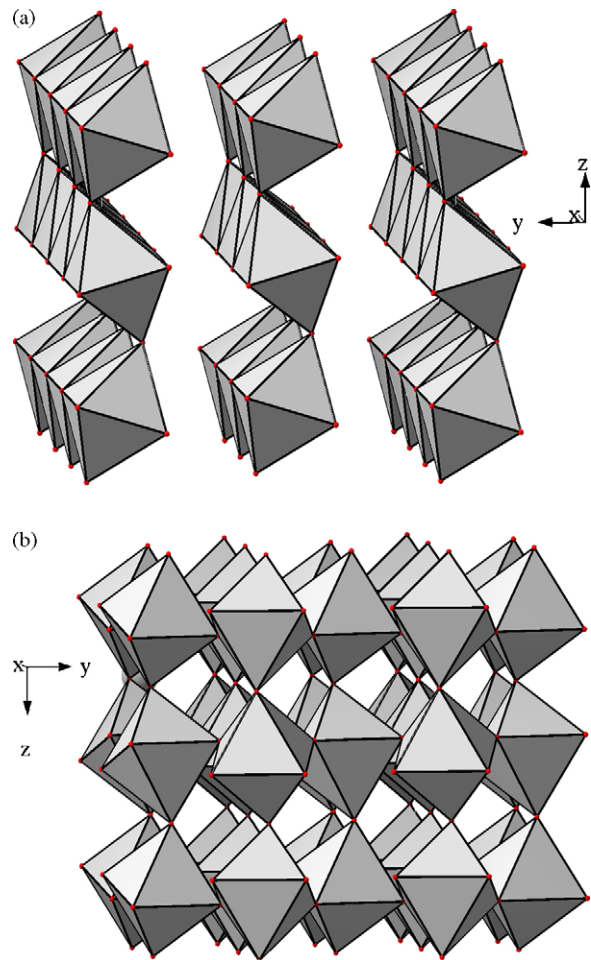


Fig. 1. Polyhedron representations of the post-perovskite (*Cmcm*) (upper) and the perovskite (*Pbnm*) (lower) CaIrO₃ crystal structures.

$5 \times 5 \times 5$ for *Pbnm*. The chosen parameters give approximately the same *k*-point density in the reciprocal space. The Monkhorst-Pack mesh was used for both cells. All calculations correspond to $T = 0$ K and $P = 0$ bar. The calculated lattice vectors are compared with experimental values in Table 1.

The magnetic moment of iridium in ferromagnetic pv and ppv were calculated since the spin-state is impor-

Table 1

Lattice vectors for post-perovskite (*Cmcm*) and perovskite (*Pbnm*) CaIrO₃ at 1 bar and 0 K

Phase	<i>a</i> (Å)	<i>b</i> (Å)	<i>c</i> (Å)	<i>V</i> (Å ³)	Reference
ppv	3.2217	9.8487	7.3502	233.22	This work
ppv	3.144	9.865	7.297	226.32	Hirose and Fujita (2005)
ppv	3.14 ₅	9.85 ₅	7.29 ₃	226.04	Rodi and Babel (1965)
pv	5.3593	5.6905	7.7341	235.87	This work
pv	5.349	5.592	7.679	229.69	Hirose and Fujita (2005)

tant for the entropy of the transition. Iridium was in both structures found to be in the low-spin state. All other calculations were performed on non-spin polarized (NSP) systems. This is in part due to the computational cost and in part due to the lack of experimental information on the magnetic ordering. This simplification is not believed to affect the results significantly. Experimental studies of magnetic order–disorder transitions for inorganic compounds generally suggest that the difference in order–disorder enthalpies between two structures of a compound is small and that the effect of magnetic order on the lattice heat capacity is negligible.

2.3. Calculation of bulk modulus and its pressure derivative from $E(V)$

The bulk modulus, K_0 , and its pressure derivative K'_0 can be deduced from the variation of the total energy with volume. In the present study, both the Murnaghan and the Birch–Murnaghan equations of state were used (Andreson, 1995). The bulk modulus values were derived by calculating the total energy at different volumes, allowing all atomic positions and all lattice vectors to vary while keeping the volume fixed. The variation of the lattice vectors and volume with pressure were subsequently deduced using the Birch–Murnaghan equation of state to determine the pressure corresponding to a given volume. The bulk moduli for pv and ppv were also calculated imposing a constant K'_0 of 4.0.

2.4. Calculation of elastic constants

The bulk modulus and the shear modulus can alternatively be determined from the elastic constants which determine the stiffness of the crystal against externally applied stresses. The elastic constants are presently determined by applying small strains to the equilibrium unit cells. For small deformations we expect Hooke's law to be valid and thus a quadratic dependence of the total

energy on the strain. This quadratic behavior is described by the elastic constants, C_{ijkl} , that are derived through a Taylor expansion of the total energy for the system $E(V, \varepsilon)$ with respect to the strain ε (Wallace, 1972):

$$E(V, \varepsilon_{mn}) = E(V_0, 0) + V_0 \left(\sum_{ij} \sigma_{ij} \varepsilon_{ij} + \frac{1}{2} \sum_{ijkl} C_{ijkl} \varepsilon_{ij} \varepsilon_{kl} \right) + O(\varepsilon^3) \quad (1)$$

V_0 is the volume of the unstrained system with a corresponding energy $E(V_0, 0)$. It is useful to use Voigt notation which takes advantage of the symmetries of the tensors: $xx \rightarrow 1, yy \rightarrow 2, zz \rightarrow 3, yz \rightarrow 4, xz \rightarrow 5, xy \rightarrow 6$. Using this notation, Eq. (4) becomes:

$$E(V, e_i) = E(V_0, 0) + V_0 \left(\sum_i \sigma_i e_i + \frac{1}{2} \sum_{ij} C_{ij} e_i e_j \right) + O(e^3) \quad (2)$$

with the strain tensor given by

$$\varepsilon = \begin{pmatrix} e_1 & \frac{1}{2}e_6 & \frac{1}{2}e_5 \\ \frac{1}{2}e_6 & e_2 & \frac{1}{2}e_4 \\ \frac{1}{2}e_5 & \frac{1}{2}e_4 & e_3 \end{pmatrix} \quad (3)$$

In order to calculate the nine elastic constants ($C_{11}, C_{22}, C_{33}, C_{12}, C_{13}, C_{23}, C_{44}, C_{55}$ and C_{66}) of the orthorhombic structure, we apply nine independent strains to the lattice vectors. The parameterizations of the nine strains used here are given in Table 2. Variable-volume distortions are used since all calculations are performed at zero pressure. For each of the strains, the total energy was calculated for five to seven distortions, γ . The zero-, first- and second-order coefficients were found through polynomial fits. By solving the nine linearly independent equations, the nine elastic constants

Table 2

Parameterization of the nine strains used to calculate the nine elastic constant of the two orthorhombic structures

Strain I	Parameters (unlisted $e_i = 0$)	$\Delta E/V$ to $O(\gamma^2)$
1	$e_1 = \gamma$	$1/2 C_{11} \gamma^2$
2	$e_2 = \gamma$	$1/2 C_{22} \gamma^2$
3	$e_3 = \gamma$	$1/2 C_{33} \gamma^2$
4	$e_1 = 2\gamma, e_2 = -\gamma, e_3 = -\gamma$	$1/2 (4C_{11} - 4C_{12} - 4C_{13} + C_{22} + 2C_{23} + C_{33}) \gamma^2$
5	$e_1 = -\gamma, e_2 = 2\gamma, e_3 = -\gamma$	$1/2 (C_{11} - 4C_{12} - 2C_{13} + 4C_{22} - 4C_{23} + C_{33}) \gamma^2$
6	$e_1 = -\gamma, e_2 = -\gamma, e_3 = 2\gamma$	$1/2 (C_{11} + 2C_{12} - 4C_{13} + C_{22} - 4C_{23} + 4C_{33}) \gamma^2$
7	$e_4 = \gamma$	$1/2 C_{44} \gamma^2$
8	$e_5 = \gamma$	$1/2 C_{55} \gamma^2$
9	$e_6 = \gamma$	$1/2 C_{66} \gamma^2$

The energy expression is given by Eq. (2).

are found. The fractional coordinates were relaxed during all calculations.

2.5. Calculation of phonon dispersion relation

The phonon dispersion relations were calculated by the direct method using the PHONON software (Parlinski, 1999; Parlinski et al., 1997). Supercells were used in order to obtain accurate frequencies at several high symmetry points of the reciprocal lattice. For ppv-CaIrO₃ a 120 atoms $3 \times 1 \times 2$ super cell was sampled with a $2 \times 2 \times 2$ k -point mesh. A 80 atom $2 \times 1 \times 2$ super cell was sampled with a $2 \times 3 \times 2$ k -point mesh for pv-CaIrO₃. The atoms at unique lattice sites were displaced one at a time by 0.03 Å along the x , y and z -directions of the unit cell and the resulting Hellmann–Feynman forces calculated. To diminish systematical errors both positive and negative displacements were made about the equilibrium positions. Twenty-four independent displacements were made for both structures. The splitting of optical modes to transverse (TO) and longitudinal (LO) components is neglected.

3. Results

3.1. Structure, energetics and compressibility

The ground state properties of ppv-CaIrO₃ and pv-CaIrO₃ were derived by relaxing both the lattice vectors and all fractional coordinates. ppv-CaIrO₃ is found to be the lower-energy structure and thus the low-temperature modification in agreement with experiments. The calculated fractional coordinates for the different ions for ppv-CaIrO₃ and pv-CaIrO₃ are given in Table 3. In case of ppv-CaIrO₃ experimental values (Rodi and Babel, 1965) are also included in the table. The calculated structural parameters given in Tables 1 and 3 are in good agreement with the experimental values. In general, GGA overestimates equilibrium volumes, and agreement with experiment within 2% is considered good. The energy difference between the two modifications was calculated to be 11.8 kJ mol⁻¹. This is somewhat lower than the recently reported calorimetric value 32.2 ± 15.5 kJ mol⁻¹ (Kojitani et al., 2007).

The bulk moduli and their pressure derivatives are deduced (using the Murnaghan and Birch–Murnaghan equations of state) from the variation of the total energy of the two structures with volume as given in Fig. 2. Only data points within $0.90 < V/V_0 < 1.10$ were used in the analyses. The bulk moduli obtained using the two different equations of state agree within 0.5%, giving $K_0 = 164$

Table 3

Fractional coordinates for post-perovskite (*Cmcm*) and perovskite (*Pbmm*) CaIrO₃ at 1 bar and 0 K

			This work	Rodi and Babel (1965)
ppv-CaIrO ₃				
Ca	4c	y	0.250	0.2498
O	4c	y	0.924	0.933
O	8f	y	0.626	0.630
O	8f	z	0.452	0.445
This work				
pv-CaIrO ₃				
Ca	4c	x		0.978
Ca	4c	y		0.064
O	4c	x		0.106
O	4c	y		0.458
O	8d	x		0.687
O	8d	y		0.303
O	8d	z		0.054

and 178 GPa for ppv and pv, respectively (Table 4). The bulk modulus of ppv-CaIrO₃ at ambient pressure is about 9% lower than that for pv-CaIrO₃. A pressure derivative, K'_0 of 3.9–4.0 for ppv-CaIrO₃ compared to 2.8–3.3 for pv-CaIrO₃ suggests that the difference in bulk modulus decreases with increasing pressure. High-pressure experiments and computations on MgSiO₃-system have yielded similar, but smaller magnitude, K'_0 -differences between pv (3.9–4.1) (Oganov and Ono, 2004; Fiquet et al., 2000) and ppv (4.2–4.5) (Oganov and Ono, 2004; Tsuchiya et al., 2005; Tsuchiya et al., 2004). Because the difference in the pressure derivatives of the bulk modulus for pv and ppv is larger in our study compared with the results from pv and ppv in the MgSiO₃-system (e.g. Oganov and Ono, 2004; Fiquet et al., 2000), we also

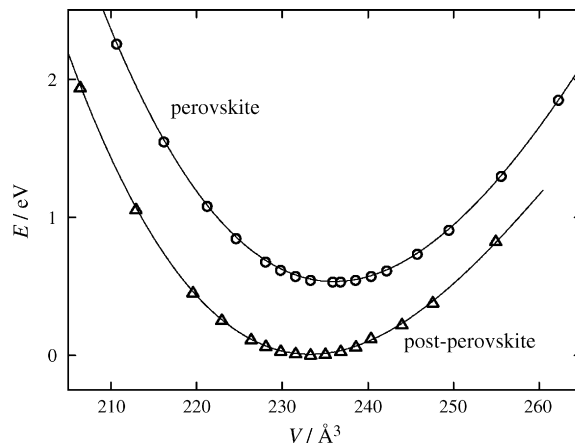


Fig. 2. E vs. V for perovskite (*Pnma*) and for post-perovskite (*Cmcm*) CaIrO₃ ($p = 1$ bar and $T = 0$ K).

Table 4
Calculated mechanical properties for post-perovskite (*Cmcm*) and perovskite (*Pbnm*) CaIrO_3 at 1 bar and 0 K

	Method	ppv	pv
K_0 (GPa)	Murnaghan	164.0	178.5
K'_0	Murnaghan	3.9	2.8
K_0 (GPa)	Murnaghan (K'_0 fixed to 4)	165.6	171.7
K_0 (GPa)	Birch–Murnaghan	164.3	177.7
K'_0	Birch–Murnaghan	4.0	3.3
K_0 (GPa)	Birch–Murnaghan (K'_0 fixed to 4)	164.1	172.4
K_0 (GPa)	Elastic constants—Voigt approximation	160.0	174.0
K_0 (GPa)	Elastic constants—Reuss approximation	153.4	168.7
G (GPa)	Elastic constants—Voigt approximation	73.8	60.6
G (GPa)	Elastic constants—Reuss approximation	64.2	56.7
E' (GPa)	Elastic constants	180.5	158.1
ν	Elastic constants	0.31	0.35

calculated the bulk modulus with a fixed value for K'_0 . The effect on the calculated bulk modulus of keeping the K'_0 -value fixed at 4.0 is, as expected, very small for ppv (with independently calculated K'_0 of 3.9–4.0). For pv, however, this reduces the bulk modulus from about 178 to 172 GPa.

There is limited previous information about the compressibility of the CaIrO_3 polymorphs. The only experimental value we have found is a pressure derivative of the bulk modulus for ppv- CaIrO_3 . The value, derived from thermal expansion data using a second-order Grüneisen approximation ($K'_0 = 4.8 \pm 0.4$) (Lindsay-Scott et al., 2007), is somewhat higher than the value obtained here. Lindsay-Scott et al. (2007) also estimated the bulk modulus for ppv- CaIrO_3 , assuming the Grüneisen parameter obtained in a DFT study of MgSiO_3 (Tsuchiya et al., 2004) to hold also for CaIrO_3 . The resulting K_0 of 144 GPa is somewhat lower than the presently derived value.

The variation of the lattice vectors and volume with pressure are deduced from the Birch–Murnaghan equation of state using the obtained bulk moduli and their pressure derivatives. The changes in unit cell volumes (V/V_0) and individual crystallographic dimensions (a/a_0 , b/b_0 , c/c_0) as a function of pressure for ppv- and pv- CaIrO_3 are shown in Figs. 3 and 4. Both crystal structures are compressed in an anisotropic

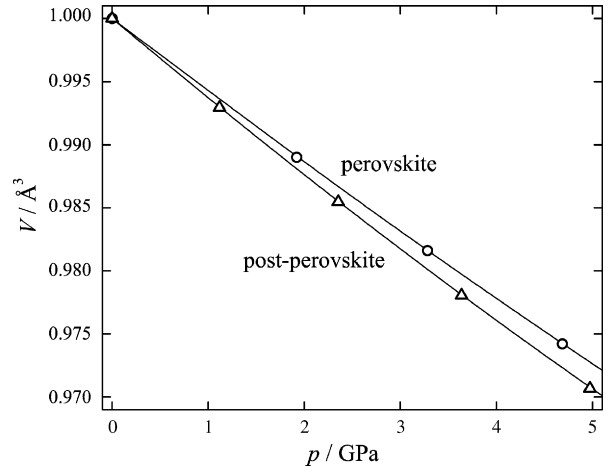


Fig. 3. V vs. p for perovskite (*Pnma*) and for post-perovskite (*Cmcm*) CaIrO_3 ($T = 0$ K, p derived from equation of state).

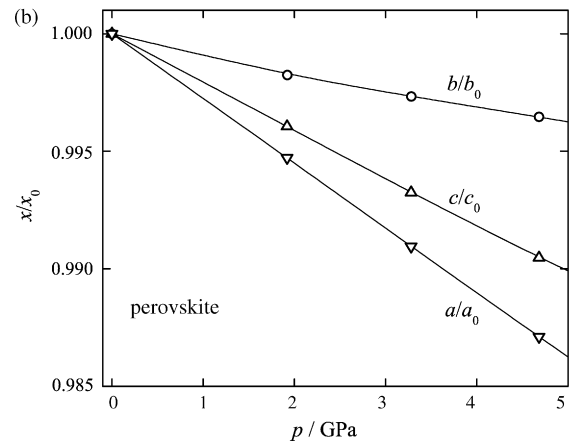
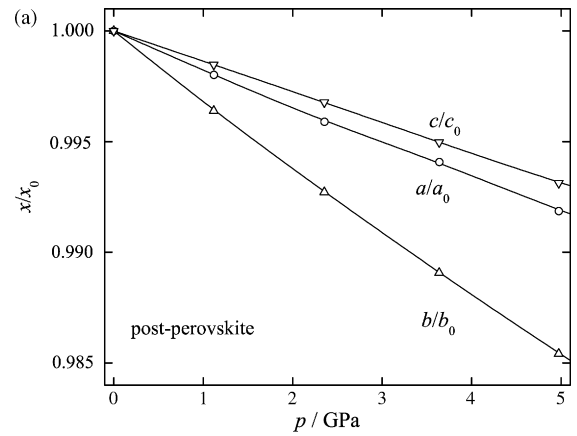


Fig. 4. (a) Lattice vectors vs. p for post-perovskite (*Cmcm*) CaIrO_3 ($T = 0$ K, p derived from equation of state), (b) Lattice vectors vs. p for perovskite (*Pnma*) CaIrO_3 ($T = 0$ K, p derived from equation of state).

fashion. For ppv-CaIrO₃ (Fig. 4a) the *a*- and *c*-axes are as expected less compressible than the *b*-axis which is perpendicular to the two-dimensional *ac*-layers.

$$K_0(R) = \frac{15}{4(S_{11} + S_{22} + S_{33}) - 4(S_{12} + S_{13} + S_{23}) + 3(S_{44} + S_{55} + S_{66})} \quad (6)$$

3.2. Elastic constants and derived mechanical properties

The elastic constants derived from the nine parameterizations of strains given in Table 2 are listed in Table 5. The constants obtained for pv- and ppv-CaIrO₃ satisfy all conditions required for mechanical stability (Wallace, 1972). The ppv-CaIrO₃ constants are clearly influenced by the layered structure resulting from edge-sharing of octahedra along the *a*-axis. This is reflected by high *C*₁₁ and *C*₃₃, compared to *C*₂₂ and low *C*₁₂ and *C*₁₃ compared to *C*₂₃ (*C*₂₂ is 61% of the average of *C*₁₁ and *C*₃₃ and the average of *C*₁₂ and *C*₁₃ is 64% of *C*₂₃). In contrast, the elastic constants for pv-CaIrO₃ are relatively similar, with *C*₁₁, *C*₃₃ and *C*₂₂ all within 21% of each other.

Polycrystalline bulk modulus and shear modulus are calculated from the elastic constants assuming a crystalline aggregate of randomly orientated single phase monocrystals. The determination of the stress or strain distribution in the aggregate with respect to external load can be derived for two extreme cases, uniform stress and uniform strain (Green, 1998). These approximations are termed the Reuss and Voigt approximations, respectively, and represent the upper and lower bounds for the elastic moduli of polycrystalline samples.

$$K_0(R) = \frac{1}{(S_{11} + S_{22} + S_{33}) + 2(S_{12} + S_{13} + S_{23})} \quad (4)$$

Table 5
Calculated elastic constants for post-perovskite (*Cmcm*) and perovskite (*Pbnm*) CaIrO₃ at 1 bar and 0 K

	ppv	pv
<i>C</i> ₁₁	325.9	224.2
<i>C</i> ₂₂	206.7	284.0
<i>C</i> ₃₃	354.9	262.8
<i>C</i> ₁₂	85.5	159.7
<i>C</i> ₁₃	69.2	103.6
<i>C</i> ₂₃	121.7	134.2
<i>C</i> ₄₄	58.6	70.4
<i>C</i> ₅₅	45.9	41.3
<i>C</i> ₆₆	61.2	66.6

$$K_0(V) = \frac{1}{9}(C_{11} + C_{22} + C_{33}) + \frac{2}{9}(C_{12} + C_{13} + C_{23}) \quad (5)$$

$$G(V) = \frac{1}{15}(C_{11} + C_{22} + C_{33} - C_{12} - C_{13} - C_{23}) + \frac{1}{5}(C_{44} + C_{55} + C_{66}) \quad (7)$$

Here *C*_{*ij*} are the elastic constants while *S*_{*ij*} are the corresponding compliance constants. Estimates of the Poisson's ratio, *ν*, and the Young's modulus, *E*', are obtained by substituting the average values of *K*₀ and *G* in:

$$\nu = \frac{1}{2} \left(1 - \frac{3G}{3K_0 + G} \right) \quad (8)$$

and

$$E' = \left(\frac{1}{3G} + \frac{1}{9K_0} \right)^{-1} \quad (9)$$

Values of the bulk and shear moduli, Poissons ratio and Youngs modulus are given in Table 4. In agreement with the results from the Birch–Murnaghan equation of state, the bulk modulus of ppv-CaIrO₃ is about 14 GPa (9%) lower than that of pv-CaIrO₃. In contrast, the shear modulus of ppv is about 10 GPa (15%) higher than that of pv.

The elastic anisotropy is characterized by the differences between the linear bulk moduli as well as the shear anisotropic factors. The bulk moduli along the *a*, *b* and *c*-axes can be calculated from the presently obtained compliance factors (Nye, 1957) as

$$K_a = \frac{1}{(S_{11} + S_{12} + S_{13})} \quad (10)$$

$$K_b = \frac{1}{(S_{12} + S_{22} + S_{23})} \quad (11)$$

$$K_c = \frac{1}{(S_{13} + S_{23} + S_{33})} \quad (12)$$

The derived values, listed in Table 6, are in very good agreement with the variation of the lattice vectors with pressure (Fig. 4) deduced from the Birch–Murnaghan equation-of-state analysis.

The shear anisotropic factors are (Ravindran et al., 1998):

$$A_1 = \frac{4C_{44}}{(C_{11} + C_{33} - 2C_{13})} \quad (13)$$

Table 6
Linear bulk moduli (in GPa) and shear anisotropy factors defined by Eqs. (13)–(15)

	CaIrO ₃		MgSiO ₃		Al ₂ O ₃	
	ppv	pv	ppv	pv	ppv	Pv
K_a	517	359	1894	1422	2294	1761
K_b	307	1122	1362	2294	2037	3205
K_c	755	444	2058	1822	2110	1721
K_b/K_a	0.59	3.12	0.72	1.61	0.89	1.82
K_c/K_a	1.46	1.24	1.09	1.28	0.92	0.98
A_1	0.43	1.01	0.72	1.19	0.64	1.07
A_2	0.58	0.59	0.92	0.89	0.55	0.80
A_3	0.68	1.49	1.06	1.27	1.04	0.98

Data for MgSiO₃ (117 GPa, 2300 K) and Al₂O₃ (136 GPa, 0 K) are taken from Wookey et al. (2005) and Stackhouse et al. (2005a,b).

For the {100} shear planes in ⟨010⟩ and ⟨011⟩ directions:

$$A_2 = \frac{4C_{55}}{(C_{22} + C_{33} - 2C_{23})} \quad (14)$$

For the {010} shear planes in ⟨001⟩ and ⟨101⟩ directions:

$$A_3 = \frac{4C_{66}}{(C_{11} + C_{22} - 2C_{12})} \quad (15)$$

For the {001} shear planes in ⟨100⟩ and ⟨110⟩ directions. The shear anisotropy factors are also given in Table 6. The deviation of the anisotropy factors from unity is a measure for the elastic anisotropy.

3.3. Lattice dynamics and vibrational properties

The calculated phonon dispersion relations for ppv- and pv-CaIrO₃ are shown in Fig. 5a and b. At the high symmetry points of the lattice a large degree of degeneration is seen. The perovskite and post-perovskite phases contain 4 and 2 formula units per primitive unit cell, respectively. According to group theory, the 57+3 phonon modes at Γ -point for the perovskite structure decompose as irreducible representation: $7A_g + 5B_{1g} + 7B_{2g} + 5B_{3g} + 8A_u + 10B_{1u} + 8B_{2u} + 10B_{3u}$. Similarly for the post-perovskite, the 27+3 phonon modes at the Γ -point decompose as $4A_g + 3B_{1g} + 1B_{2g} + 4B_{3g} + 2A_u + 6B_{1u} + 6B_{2u} + 4B_{3u}$. For both structures, all g modes are Raman active, while all u modes except the A_u ones are IR active. Further details will be published separately.

The normalized phonon densities of states, $g(\omega)$, for the supercells shown in Fig. 6 are found by integration. A simple harmonic approximation is used to calculate the molar heat capacity from these phonon densities of

states, $g(\omega)$. The heat capacity is per formula unit:

$$C_V = nk_B \int_0^\infty \left(\frac{\hbar\omega}{kT} \right) \frac{e^{\hbar\omega/k_B T}}{(e^{\hbar\omega/k_B T} - 1)^2} g(\omega) d\omega \quad (16)$$

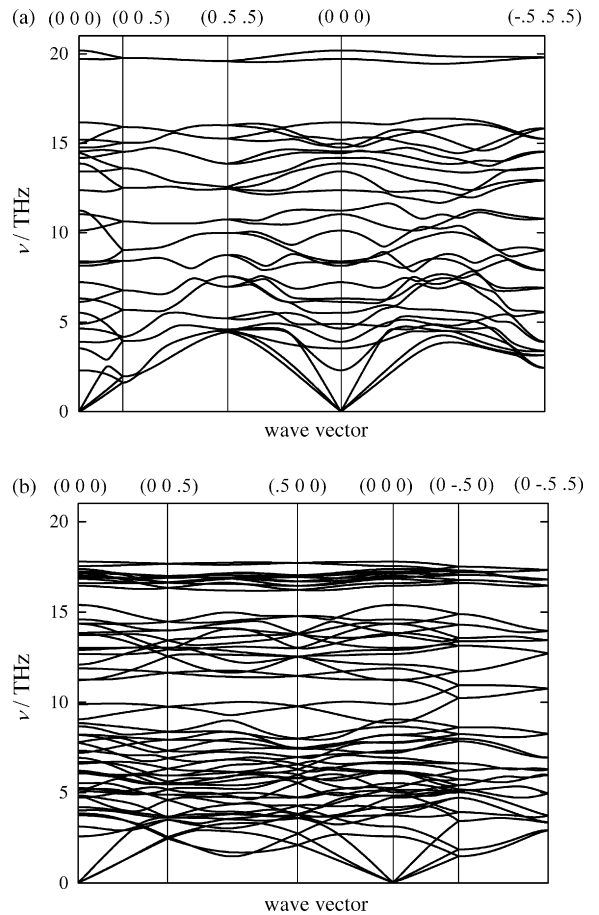


Fig. 5. (a) Dispersion relation for post-perovskite (*Cmcm*) CaIrO₃ ($p=1$ bar and $T=0$ K). (b) Dispersion relation for perovskite (*Pnma*) CaIrO₃ ($p=1$ bar and $T=0$ K).

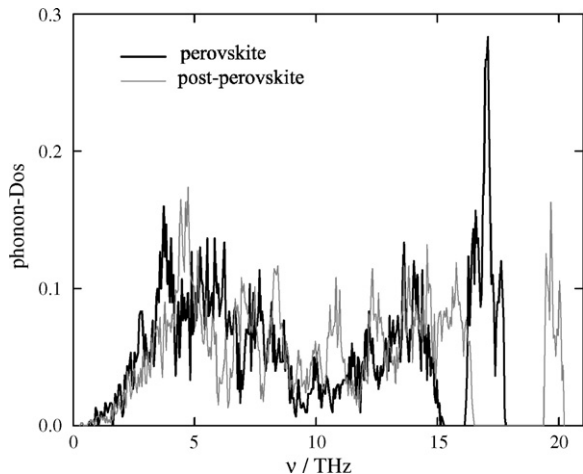


Fig. 6. Phonon density of state for perovskite (*Pnma*) and for post-perovskite (*Cmcm*) CaIrO_3 ($p = 1$ bar and $T = 0$ K).

where n is the number of atoms, \hbar the Planck constant and k_B is the Boltzmann constant. The calculated molar heat capacities and the derived molar entropies are given in Figs. 7 and 8, respectively. At 298 K the difference in entropy is $6.3 \text{ J K}^{-1} \text{ mol}^{-1}$. Above 298 K the heat capacities of the two phases are very similar and the entropy difference increases only by $0.4 \text{ J K}^{-1} \text{ mol}^{-1}$ between 298 and 1500 K. A main question is to what extent these entropies converge with regards to cell-size. We have addressed that question by calculating the thermodynamic parameters also for the $1 \times 1 \times 1$ unit cells which give exact vibrational frequencies only at the Brillouin zone center (the Γ -point). The molar entropy at 298 K increases from 103.0 to $105.5 \text{ J K}^{-1} \text{ mol}^{-1}$ going from a $1 \times 1 \times 1$ unit cell to a $2 \times 1 \times 2$ super cell for pv- CaIrO_3 . For ppv- CaIrO_3 ,

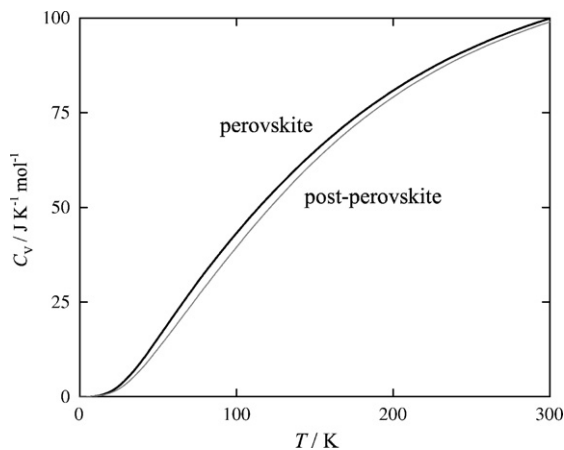


Fig. 7. Molar heat capacity for perovskite (*Pnma*) and for post-perovskite (*Cmcm*) CaIrO_3 at 1 bar.

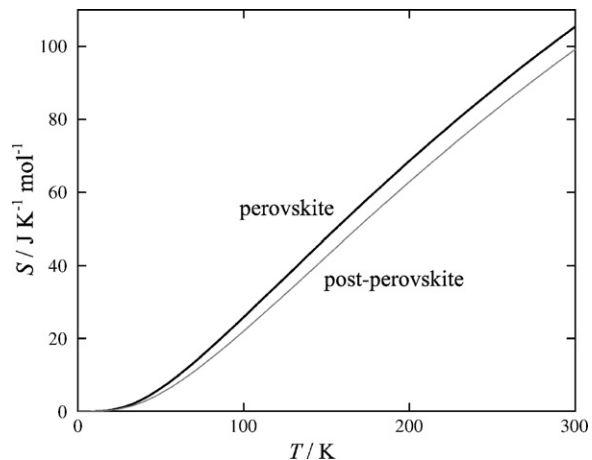


Fig. 8. Molar entropy for perovskite (*Pnma*) and for post-perovskite (*Cmcm*) CaIrO_3 at 1 bar.

the molar entropy decreases from $99.8 \text{ J K}^{-1} \text{ mol}^{-1}$ for a $1 \times 1 \times 1$ unit cell to $99.2 \text{ J K}^{-1} \text{ mol}^{-1}$ for a $3 \times 1 \times 2$ super cell. The entropy difference thus increased from 3.2 to $6.3 \text{ J K}^{-1} \text{ mol}^{-1}$ at 298 K. The change in entropy is expected to be small when increasing the size of the super cell further. This expectation, however, cannot be tested due to the prohibitive cost involved.

3.4. High pressure behaviour

Based on the calculations presented above we can estimate the slope of the ppv–pv phase boundary. The Clapeyron equation may be written as

$$\frac{dp}{dT} = \frac{\Delta S}{\Delta V} = \frac{\Delta H}{T_{\text{trs}} \Delta V} \quad (17)$$

We are not able to produce a pure *ab initio* value of the slope since we are not able to calculate all contributions to the entropy of transition. However, we can calculate a slope based on the calculated enthalpy of transition (11.8 kJ mol^{-1}) and the calculated volume of transition (0.64 \AA^3 per formula unit or $0.40 \text{ cm}^3 \text{ mol}^{-1}$). By assuming an extrapolated experimental value for a hypothetical transition temperature at ambient pressure of 1380°C (1673 K), we obtain a dp/dT -slope of 18 MPa K^{-1} . The resulting phase boundary is compared with literature values in Fig. 9. It is difficult to estimate the accuracy of this value and it should be noticed that the calculated and experimental volumes of transition varies between 0.32 and $0.51 \text{ cm}^3 \text{ mol}^{-1}$ (Hirose and Fujita, 2005; Trønnes et al., 2006). This dp/dT -slope is intermediate between the experimentally determined slopes of about 16 MPa K^{-1} (Hirose and Fujita, 2005) and 24 MPa K^{-1} (Trønnes et al., 2006). In a solution calorimetric study Kojitani et al.

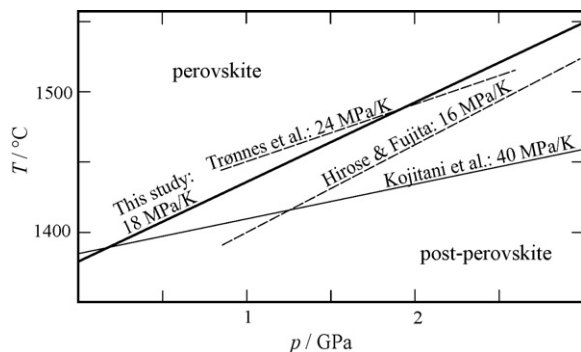


Fig. 9. Compilation of pv–ppv phase boundaries for CaIrO_3 from various studies. Phase equilibria from quenching experiments: Hirose and Fujita (2005) and Trønnes et al. (2006); from calorimetry: Kojitani et al. (2007); from DFT-modeling: this study.

(2007) obtained a very large, and rather uncertain slope of $40 \pm 19 \text{ MPa K}^{-1}$, with the quoted uncertainty based on the standard deviations reported for their enthalpy change.

4. Discussion

Our computationally derived enthalpy difference between the two CaIrO_3 polymorphs is small compared to the only experimental value (Kojitani et al., 2007), and further experimental and computational studies are needed to accurately resolve this issue. The solution calorimetric determination of the enthalpy difference between pv and ppv- CaIrO_3 (Kojitani et al., 2007) suggest that the entropy of ppv is as much as $19.4 \pm 9.3 \text{ J K}^{-1} \text{ mol}^{-1}$ lower than for pv. Kojitani et al. (2007) suggest the difference in structural connectivity as the origin of the entropy difference. Whereas pv- CaIrO_3 consists of corner-sharing octahedra only, the ppv-structure contains edge-sharing octahedra along the a -axis. The more rigid edge-linking may restrict some phonons, in particular those involving rotational motion of the octahedra (Kojitani et al., 2007). Although we in principle agree with this line of reasoning, the magnitude of this effect is probably insufficient to account for the large entropy difference obtained by (Kojitani et al., 2007). In a recent study of SrMnO_3 (Søndenå et al., 2007) we determined, both experimentally and computationally, the heat capacity and entropy of two polymorphs differing even more in connectivity. Whereas cubic SrMnO_3 has corner-sharing octahedra only, the hexagonal polymorph contains 50% corner-sharing and 50% face-sharing octahedra. In this case the entropy change at 298 K is as small as $5.5 \text{ J K}^{-1} \text{ mol}^{-1}$ (Søndenå et al., 2007). We do not expect the difference in the dilational contribution to the heat capacity, and thus to the entropy,

to be large enough to explain the difference between the entropy change based on the solution calorimetry (Kojitani et al., 2007) and that obtained here. On the other hand, it is clear that electronic contributions to the heat capacity and thus the entropy are significant if pv- or pv- CaIrO_3 or both are metallic. Information on the electrical properties of CaIrO_3 , however, are not available. Other ternary iridium oxides such as BaIrO_3 and Sr_2IrO_4 in which the Ir is in the formal Ir^{4+} -state are known to show weak ferromagnetism with a relatively high disordering temperature and to show unusual electrical properties with overall semiconducting behavior (Kini et al., 2006). It is important to note that both the pv and ppv modifications are low-spin at least in the ferromagnetic state. A difference in spin state between the two structures could have induced a significant entropy of transition. Thus, the entropy difference between pv and ppv appears to be smaller than that proposed by Kojitani et al. (2007). The standard deviation reported in that study, however, is very large.

The entropy difference between pv- and ppv- CaIrO_3 can also be estimated from the experimental determination of the Clapeyron slopes and the volume change of the transition (Hirose and Fujita, 2005; Trønnes et al., 2006). The values of 16 MPa K^{-1} and $0.507 \text{ cm}^3 \text{ mol}^{-1}$ (Hirose and Fujita, 2005) and 24 MPa K^{-1} and $0.321 \text{ cm}^3 \text{ mol}^{-1}$ (Trønnes et al., 2006) gives entropy differences between pv and ppv of 8.11 and $8.02 \text{ J K}^{-1} \text{ mol}^{-1}$, respectively. These entropy changes are in reasonable agreement with the entropy difference estimated presently from the calculated phonon dispersion relations, $6.3 \text{ J K}^{-1} \text{ mol}^{-1}$. In conclusion, the present study and the previously determined Clapeyron slopes give a consistent picture that indicate that the enthalpy and thus the entropy of the pv to ppv transition is overestimated in the recent calorimetric study (Kojitani et al., 2007).

The negative and positive changes in bulk and shear moduli, respectively, across the pv- to ppv-transition boundary is strongly linked to the nature of the crystal structure. Edge-sharing octahedra along the a -axis in the ppv-structure give rise to high values of C_{11} and C_{33} , i.e. high resistance towards stress in the ac -plane of the 2D slabs. For the same reason C_{12} and C_{13} are low for the ppv-structure. Whereas the bulk modulus (in the Voigt approximation) is related to the sum of $(C_{11} + C_{22} + C_{33})$ and $(C_{12} + C_{13} + C_{23})$, the shear modulus is related to their difference. The bulk and shear moduli changes across the transition is thus explained by the higher $(C_{11} + C_{22} + C_{33})$ - and lower $(C_{12} + C_{13} + C_{23})$ -sum of the ppv- compared to the pv-structure.

Table 6 presents a comparison of the elastic anisotropy of CaIrO_3 , MgSiO_3 and Al_2O_3 based on this study, (Wookey et al., 2005) and (Stackhouse et al., 2005a,b), respectively. It should be noted that Oganov and Ono (2005) concluded that a $\text{Rh}_2\text{O}_3(\text{II})$ -structured phase of Al_2O_3 transforms directly to ppv- Al_2O_3 at about 120 GPa and 0–2500 K, without any intervening pv-stability field. They indicated, however, that there may be a pv-stability field at temperatures above 2500 K. With this caveat for Al_2O_3 in mind, it is interesting to note the close correspondence between the various linear bulk moduli and shear anisotropy factors of pv and ppv for the three different compositions. The similar elastic properties for the systems CaIrO_3 , MgSiO_3 and Al_2O_3 seem promising in terms of using pv- and ppv- CaIrO_3 as low-pressure analogues for the corresponding deep Earth minerals in $(\text{Mg},\text{Fe}^{2+})\text{SiO}_3$ -dominated systems with minor, but significant Al_2O_3 - and Fe_2O_3 -components.

A comparison between the compressional and thermal anisotropy of ppv- CaIrO_3 can be based on our DFT-computational study and the experimental determination of thermal expansivity by Lindsay-Scott et al. (2007). These results indicate that the pressure-induced elastic anisotropy is considerably larger than the thermal anisotropy. Whereas the *b*-axis has the greatest thermal expansion and isothermal compressibility, the least compressible *c*-axis does not correspond to the direction of the lowest thermal expansion, i.e. the *a*-axis (Lindsay-Scott et al., 2007). The strong elastic anisotropy of ppv, documented from DFT-computations of both MgSiO_3 and CaIrO_3 corresponds to the well-developed seismic anisotropy of D'' -zone (Lay et al., 2006; Wookey et al., 2005).

The *Pbnm* perovskites of CaIrO_3 and MgSiO_3 have open and flexible crystal structures with corner-linked octahedra and high entropy. The phase transitions into and out of the perovskite stability field are therefore endothermic and exothermic, respectively. The moderately endothermic perovskite-forming reaction in MgSiO_3 -dominated compositions associated with the 660 km seismic discontinuity has a negative dp/dT -slope of -2 to -3 MPa K^{-1} (Weidner and Wang, 2000), whereas the exothermic post-perovskite-forming transitions defining the main D'' -discontinuity have a much larger and positive dp/dT -slope. The exothermic nature of the post-perovskite forming reaction is enhanced by the notably low entropy of post-perovskite, associated with a major structural reorganization and transition to 33% corner-linked octahedra. The 660 km discontinuity has a relatively simple seismological expression in the form of increasing v_p and v_s . The sharp discontinuity

is followed by velocity gradients in the 660–720 km depth range associated with delayed transformation of majoritic garnet to more aluminous perovskite (Kennett et al., 1995; Shearer, 2000). The D'' discontinuities, however, are considerably more complicated. As indicated by Lay et al. (2006), the uppermost discontinuity, associated with a minor velocity decrease, may be caused by densification associated with compositional layering. The next deeper and major discontinuity, characterized by increasing v_s and largely unchanged or even slightly decreasing v_p , requires the unusual combination of increasing shear modulus and decreasing bulk modulus across the interphase. A lower and reverse discontinuity (decreasing v_s) within 50–100 km of the core–mantle boundary (Lay et al., 2006) may represent a transition where post-perovskite reverts to perovskite due to large thermal gradients near the outer core.

The shear and bulk moduli requirements for the two latter discontinuities are consistent with the DFT-computational results of Wookey et al. (2005) for MgSiO_3 and this study for the CaIrO_3 -analogue composition. Other *ab initio* calculations at 100–120 GPa (Oganov and Ono, 2004; Itaka et al., 2004) show slightly smaller increases of shear modulus across the MgSiO_3 pv to ppv phase transition (3–5% versus 6–8% in (Wookey et al., 2005)) and almost no change in the bulk modulus. High-pressure experimental bulk modulus determinations for MgSiO_3 , however, show decreases across the pv–ppv transition of 10–20% (Shieh et al., 2006; Stackhouse et al., 2005a,b), largely consistent with our computational CaIrO_3 -results.

The large Clayperon slope resulting from the small volume increase and large entropy decrease of the pv–ppv-transition has important geodynamic implications. The large topography of the discontinuity associated with lateral temperature variation near the core–mantle boundary will in principle promote convection and heat transfer in the D'' -zone. However, this effect may be largely cancelled by pronounced compositional layering, involving intrinsically dense thermochemical piles enriched in subducted basaltic material or material with elevated Fe/Mg-ratio. The observations of Torsvik et al. (2006) may indicate that sub-Pacific and sub-African thermochemical piles above the core–mantle boundary have been intact and stable over more than 250 Ma.

5. Conclusions

Our *ab initio* computational study of the thermodynamics and elasticity of the pv- and ppv-phases in the system CaIrO_3 are broadly consistent with

experiments and published DFT-results on the MgSiO_3 -composition. Therefore, the CaIrO_3 system may be a promising low-pressure analogue for the Al- and Fe-bearing MgSiO_3 -dominated pv and ppv of the lowermost mantle. The results show that the pv–ppv transition in CaIrO_3 is associated with decreasing volume (about 1%) and entropy (about 20%), resulting in a large dp/dT -slope of 18 MPa K^{-1} for the phase boundary, broadly consistent with recent experimental phase equilibrium investigations.

The calculated elastic anisotropy of pv and ppv in Ca corresponds well with the relations derived for the MgSiO_3 -polymorphs by similar studies. As expected, the b -axis of ppv (normal to the partly edge-linked octahedral networks of the ac -layers) is considerably more compressible than the a - and c -axes. In perovskite, the b -axis is the least compressible. The shear modulus increases by about 15% and the bulk modulus decreases by about 9% across the pv–ppv transition. This finding is consistent with DFT-results in the MgSiO_3 system and with seismic observation of anti-correlated v_s and v_ϕ across the main D'' -discontinuity.

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References

- Andreson, O.L., 1995. Equation of State for Solids for Geophysics and Ceramic Science. Oxford University Press.
- Fiquet, G., Dewale, A., Andrault, D., Kunz, M., Le Bihan, T., 2000. Thermoelastic properties and crystal structure of MgSiO_3 perovskite at lower mantle pressure and temperature conditions. *Geophys. Res. Lett.* 27, 21–24.
- Green, D.J., 1998. An Introduction to the Mechanical Properties of Ceramics. Cambridge University Press.
- Hernlund, J.W., Thomas, C., Tackley, P.J., 2004. A doubling of the post-perovskite phase boundary and structure of the Earth's lowermost mantle. *Nature* 434, 882–886.
- Hirose, K., Fujita, Y., 2005. Clapeyron slope of the post-perovskite phase transition in CaIrO_3 . *Geophys. Res. Lett.* 32, L13313, doi:10.1029/2005GL023219.
- Hirose, K., Takafuji, N., Sata, N., Ohishi, Y., 2005. Phase transition and density of subducted MORB crust in the lower mantle. *Earth Planet. Sci. Lett.* 237, 239–251.
- Iitaka, T., Hirose, K., Kawamura, K., Murakami, M., 2004. The elasticity of the MgSiO_3 post-perovskite phase in the Earth's lowermost mantle. *Nature* 430, 442–445.
- Kennett, B.L.N., Engdahl, E.R., Buland, R., 1995. Constraints on the velocity structure of the Earth from travel times. *Geophys. J. Int.* 122, 108–124.
- Kini, N.S., Strydom, A.M., Jeevan, H.S., Geibel, C., Ramakrishnan, S., 2006. Transport and thermal properties of weakly ferromagnetic Sr_2IrO_4 . *J. Phys.: Condens. Matter* 18, 8205–8216.
- Kojitani, H., Furukawa, A., Akaogi, M., 2007. Thermochemistry and high-pressure equilibria of the post-perovskite phase transition in CaIrO_3 . *Am. Mineral.* 92, 229–232.
- Kresse, G., Furthmüller, J., 1999. Efficient iterative schemes for *ab initio* total-energy calculations using a plane-wave basis set. *Phys. Rev. B* 54, 11169–11186.
- Kresse, G., Hafner, J., 1993. *Ab initio* molecular dynamics for liquid metals. *Phys. Rev. B* 47, 558–561.
- Lay, T., Hernlund, J., Garnero, E., Thorne, M.S., 2006. A post-perovskite lens and D'' heat flux beneath the central Pacific. *Science* 314, 1272–1276.
- Lindsay-Scott, A., Wood, I.G., Dobson, D.P., 2007. Thermal expansion of CaIrO_3 determined by X-ray powder diffraction. *Phys. Earth Planet. Int.* 162, 140–148.
- Mao, W.L., Shen, G., Prakapenka, V.B., Meng, Y., Cambell, A.J., Heinz, D., Shu, J., Hemley, R.J., Mao, H.K., 2004. Ferromagnesian postperovskite silicates in the D'' layer of the earth. *Proc. Natl. Acad. Sci. USA* 101, 15,867–15,869.
- McDaniel, C.L., Schneider, S.J., 1972. Phase relations in the $\text{CaO-} \text{IrO}_2\text{-Ir}$ system in air. *J. Solid State Chem.* 4, 275–280.
- Murakami, M., Hirose, K., Kawamura, K., Sata, N., Ohishi, Y., 2004. Post-perovskite phase transition in MgSiO_3 . *Science* 304, 855–858.
- Murakami, M., Hirose, K., Sata, N., Ohishi, Y., 2005. Post-perovskite phase transition and mineral chemistry in the pyrolytic lower mantle. *Geophys. Res. Lett.* 32, L03304, doi:10.1029/2004GL021956.
- Nakagawa, T., Tackley, P.J., 2004. Effects of a perovskite-post perovskite phase change near core-mantle boundary in compressible mantle convection. *Geophys. Res. Lett.* 31, L16611, doi:10.1029/2004GL020648.
- Nye, J.F., 1957. Physical Properties of Crystals. Clarendon Press, Oxford.
- Oganov, A.R., Ono, S., 2004. Theoretical and experimental evidence for a post-perovskite phase of MgSiO_3 in Earth's D'' layer. *Nature* 430, 445–448.
- Oganov, A.R., Ono, S., 2005. The high-pressure phase of alumina and implications for the Earth's D'' layer. *Proc. Nat. Acad. Sci.* 102, 10828–10831.
- Ono, S., Kikegawa, T., Ohishi, Y., 2006. Equation of state of the CaIrO_3 -type MgSiO_3 up to 144 GPa. *Am. Mineral.* 91, 475–478.
- Parlinski, K., 1999. Calculation of phonon dispersion curves by the direct method, *Amer. Inst. Phys., Conference Proceedings* 476 “Neutrons and Numerical Methods N.2M”, ed. M.R. Johnson, G.G. Kearley, H.G. Buttner, p. 121.
- Parlinski, K., Li, Z.Q., Kawazoe, Y., 1997. First-principles determination of the soft mode in cubic ZrO_2 . *Phys. Rev. Lett.* 78, 4063–4066.
- Perdew, J.P., Burke, K., Ernzerhof, M., 1996. Generalized gradient approximation made simple. *Phys. Rev. Lett.* 77, 3865–3868.
- Ravindran, P., Fast, L., Kozhavyi, P.A., Johansson, B., Wills, J., Eriksson, O., 1998. Density functional theory for calculation of orthorhombic crystals; Application to TiSi_2 . *J. Appl. Phys.* 84, 4891–4904.

- Rodi, F., Babel, D., 1965. Ternare oxide der übergangsmetalle. 4. Erdalkaliiridium(4)-oxide—kristallstruktur von CaIrO_3 . *Z. Anorg. Allg. Chem.* 336, 17–23.
- Shearer, P.M., 2000. Upper mantle seismic discontinuities, in: S.-i. Karato et al. (eds.) *Earth's Deep Interior, Mineral Physics and Tomography From the Atomic to the Global Scale*. Am. Geophys. Union. Monograph 117, 115–131.
- Shieh, S.R., Duffy, T.S., Kubo, A., Prakapenka, V.B., Sata, N., Hirose, K., Ohishi, Y., 2006. Equation of state of the postperovskite phase synthesized from a natural $(\text{Mg,Fe})\text{SiO}_3$ orthopyroxene. *Proc. Natl. Acad. Sci.* 103, 3039–3043.
- Sidorin, I., Gurnis, M., Helmlinger, D.V., 1999. Evidence for a ubiquitous seismic discontinuity at the base of the mantle. *Science* 286, 1326–1329.
- Søndenå, R., Stølen, S., Ravindran, P., Grande, T., 2007. Heat capacity and lattice dynamics of cubic and hexagonal SrMnO_3 : calorimetry and DFT simulations. *Phys. Rev. B.*, in press.
- Stackhouse, S., Brodholt, J.P., Price, G.D., 2005a. High temperature elastic anisotropy of the perovskite and post-perovskite polymorphs of Al_2O_3 . *Geophys. Res. Lett.* 32, L13305.
- Stackhouse, S., Brodholt, J.P., Wookey, J., Kendall, J.-M., Price, G.D., 2005b. The effect of temperature on the seismic anisotropy of the perovskite and post-perovskite polymorphs of MgSiO_3 . *Earth. Planet. Sci. Lett.* 230, 1–10.
- Torsvik, T.H., Smethurst, M.A., Burke, K., Steinberger, B., 2006. Large igneous provinces generated from the margins of the large low-velocity provinces in the deep mantle. *Geophys. J. Int.*, 1–14, doi:10.1111/j. 1365-246X.
- Trønnes, R.G., Frost, D.J., Boffa-Ballaran, T., Stølen, S., 2006. The perovskite to post perovskite transition in CaIrO_3 . American Geophysical Union Fall Meeting, MR11A97.
- Tsuchiya, T., Tsuchiya, J., Umamoto, K., Wentzcovitch, R.M., 2004. Phase transition in MgSiO_3 perovskite in the earth's lower mantle. *Earth. Planet. Sci. Lett.* 224, 241–248.
- Tsuchiya, J., Tsuchiya, J., Wentzcovitch, R.M., 2005. Vibrational and thermodynamic properties of MgSiO_3 postperovskite. *J. Geophys. Res.* 110, B02204.
- Wallace, D.C., 1972. *Thermodynamics of Crystals*. Wiley, New York.
- Weidner, D.J., Wang, Y., 2000. Phase transitions: implications for mantle structure, in: S.- i. Karato et al. (eds.) *Earth's Deep Interior, Mineral Physics and Tomography From the Atomic to the Global Scale*. Am. Geophys. Union, Monogr. 117, 215–235.
- Wookey, J., Stackhouse, S., Kendall, J.M., Brodholt, J., Price, D., 2005. Efficacy of the post-perovskite phase as an explanation for the lowermost-mantle seismic properties. *Nature* 438, 1004–1007.